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ORIGINAL ARTICLE

# Increase of order in seismic processes around large reservoir induced by water level periodic variation

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Abstract The importance of small periodic influences 8 on the complex systems behavior is well acknowledged. In the present research, the possible impact of 10 water level variation in large reservoir on the dynam-11 ics of local seismic activity was investigated. Large 12 reservoirs located in the seismically active zones are 13 often considered as a factor, quantitatively and qualita-14 tively influencing earthquakes generation. During im-15 poundment or after it, both the number and magnitude 16 of earthquakes around reservoir significantly increases. 17 After several years, these changes in earthquake gen-18 eration, named as reservoir-induced seismicity (RIS) 19 essentially decrease down to the level, when lesser earthquakes occur with lower magnitudes. To explain 21 this decrease, the authors of the present paper recently 22 proposed the model of phase synchronization of local 23 seismic activity by the periodic variation of the water 24 level - reservoir-induced synchronization of seismicity 25 (RISS). 26 Generally, RISS presumes a kind of control of local seismic activity by synchronizing small external peri-28 odic influence and hence increase of order in dynamics 29

<sup>30</sup> of regional seismic activity. To reveal these changes

J. Peinke Institute of Physics, University of Oldenburg, D-26111 Oldenburg, Germany in dynamics of phase-synchronized seismic activity around large reservoir field, seismic and water level variation data were analyzed in the present work. Laboratory stick–slip acoustic emission data as a model of natural seismicity were also analyzed.

The evidence is presented that increase of order in dynamics of daily earthquake occurrence, earthquakes temporal, and energy distribution took place around Enguri high dam water reservoir (Western Georgia) during the periodic variation of the water level in the lake.

**Keywords** Acoustic emission · Dynamics · Earthquakes · Reservoir · Water level variations

## **1** Introduction

Dynamics of different processes in natural systems 45 often are affected by complex interaction with each 46 other. In this respect, it is most important that they 47 often respond to the small external influence in sense 48 of phase synchronization even when processes are 49 poorly correlated [18]. Phase synchronization of 50 complex systems to external driving has been observed 51 in many biological systems, numerical models, and 52 laboratory experiments [14, 18, 19]. At the same 53 time, the phase synchronization effects in complex 54 environmental processes rarely are considered from 55 the quantitative dynamical analysis point of view. 56 In the present research, the character of dynamical 57 changes in the regional seismic activity under phase 58

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synchronizing forcing by water level variation in a 59 large reservoir was investigated. 60

Generally, the scientific and practical importance 61 of investigation of possible mechanisms related to the 62 dynamics of influence of high dam water reservoirs 63 64 on local earthquakes generation is well acknowledged [1, 15, 22, 23, 25, 26]. Since the mid of the past cen-65 tury, the RIS has been observed at many reservoirs and their geological, hydrological, and physical features re-67 main subjects of intense investigation.

At present, many aspects of dynamics of RIS become understandable [1, 26]. Namely, the increase of 70 seismic activity is related to the complex impact of the 71 water lake on the earth crust under the reservoir and is 72 explained by the changes in ambient stress conditions 73 due to the load (unload) of the water or, respectively, 74 to the increase of interstitial pore pressure in the rock 75 matrix beneath the reservoir due to downward perco-76 lation of fluid. It is also known that the water reservoir 77 related changes in the seismicity of surrounding area 78 - RIS, decreases after several years down to the level 79 when even lesser earthquakes may occur with lower 80 magnitudes compared to the basic level of the local 81 seismic activity [1, 15]. At the same time, the under-82 lying dynamics of the mentioned decrease of seismic 83 energy release, following the initial RIS activity is still 84 an open problem. 85 In our previous research based on field and labo-

ratory data, the evidence was presented that the de-87 crease of seismic energy release associated with RIS 88 may be caused by the periodic variation of the wa-89 ter level in the large reservoir [17]. Based on the re-90 sults of water level variation, as well as seismic and 91 stick-slip acoustic emission data analysis, it was shown 92 that small (compared to tectonic strain) periodic influ-93 ence on a complex seismic process may invoke local 94 reservoir-induced phase synchronization of seismicity. 95 As far as we deal with the weakest form of synchrony – phase synchronization [16] and taking into account the 97 earlier-mentioned lack of appropriate research, in the 98 present work we aimed to investigate the character of 99 dynamical changes in regional seismic activity caused 100 by the influence of water level periodic variation in the 101 large reservoir. 102

#### 2 Used data and methods 103

Data sets of daily water level variation have been col-104 lected at one of the largest in Europe, 272 m height, 105

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Enguri high dam reservoir located in the Western Geor-106 gia, Caucasus (42.030 N, 42.775 E) in 1973-1995 107 (Fig. 1(a)). 108

The size of the area around Enguri high dam, which can be considered sensitive to the reservoir influence, was evaluated based on the concept of energy release acceleration in the seismically critical regions. Namely, the minima of curvature parameter C (defined as C =power-law fit RMS error/linear fit RMS error) deduced from the Benioff strain  $E(t) = \sum_{i=1}^{N(t)} E_i(t)^{1/2}$  [3] was calculated. Here,  $E_i$  is the energy of the *i*th event. Location of the Enguri high dam reservoir was assumed as a "virtual epicenter of impending strong earthquake" (for details see [17], where it is shown that the radius of an area around Enguri high dam, sensitive to the reservoir influence is about 90 km). Data sets of daily occurred number of earthquakes and released daily seismic energy by seismic events above representative magnitude threshold  $M \ge 1.6$  within this 90 km area for 1973– 1995 are shown in Fig. 1(b) and (c). Besides these daily data, time series of sequences of magnitudes and time intervals between consecutive earthquakes (wait-127 ing times), unevenly sampled for the same time period 128 and area, also were analyzed. 129

Used data sets of water level variation and seismic activity around Enguri high dam were obtained from data bases of the M. Nodia Institute of Geophysics 132 (Georgia).

Acoustic emission data were collected on the stick 134 slip laboratory setup represented by a system of two 135 roughly finished basalt plates [5, 6]. The external faces 136 of plates were subjected to periodic electric (48 Hz) per-137 turbations (with amplitudes, varying from 0 to 1000 V) 138 superimposed on the constant dragging force (normal-139 ized power of an external sinusoidal forcing is shown 140 in Fig. 2(a)). The waveforms of both acoustic emission 141 and the sinusoidal EM field were digitized at 44 kHz. 142 From the digitized waveforms of acoustic emission 143 data sets, the time series (catalogs) of power of emitted 144 acoustic energy were compiled (Fig. 2(b)). Exactly, the 145 power of emitted acoustic energy was calculated as the 146 area between the acoustic signal curve and the x-axis 147 during the period of the superimposed external 48 Hz 148 sinusoidal forcing divided by the time duration of these 149  $2\pi$  periods. Additionally, sequences of time intervals 150 between consecutive maximal amplitudes of acoustic 151 signals (waiting times) were analyzed. 152

Besides formerly described in [17] characteris-153 tics, the mean effective phase diffusion coefficient 154

Fig. 1 (a) Record of the daily water level in the lake of Enguri dam from 1975 to 1993, (b) daily number of earthquakes, and (c) log of normalized daily released seismic energy



<sup>155</sup>  $D = \frac{d}{dt} [\langle \Delta \varphi^2 \rangle - \langle \Delta \varphi \rangle^2]$  was calculated as an additional statistical measure of the quality of synchronization between water level variation and seismicity, as well as between external periodic forcing and power of acoustic emission.

<sup>160</sup> In order to investigate dynamical changes in <sup>161</sup> analysed processes, Recurrence Quantitative Analysis or (RQA) was used [27]. RQA is especially useful to overcome the difficulties often encountered dealing with nonstationary and rather short real data sets. The recurrence plots (RP) are defined as: 166

$$R_{i,j} = \Theta(\varepsilon_i - \|\bar{x}_i - \mathbf{x}_j\|)$$

Fig. 2 Results of experiments on stick–slip, which is considered as a good laboratory model for seismicity, affected by weak external perturbations: (a) Normalized power of the external sinusoidal forcing,  $P_s$ , superimposed on the constant drag force; (b) normalized power of acoustic emission of stick–slip events,  $P_a$ 



Here,  $\varepsilon_i$  is a cutoff distance and  $\Theta(x)$  is the Heaviside 167 function. Correct choice of cutoff distance  $\varepsilon$  is one of 168 the main problems of RQA. It is desirable that  $\varepsilon$  be 169 as small as possible, but the presence of noise always 170 necessitates larger values. There are several sugges-171 tions how to set correctly  $\varepsilon$  [7, 12, 27]. We selected 172 the cutoff distance as 10% (for waiting times and daily 173 number of earthquakes) and 20% (for magnitude se-174 quence) of overall mean distance [2, 12]. As a quanti-175 tative tool of complex dynamics analysis, RQA defines 176 several measures mostly based on diagonally oriented 177 lines in the recurrence plots: recurrence rate, determin-178 ism, maximal length of diagonal structures, entropy, 179 trend, etc. [7]. In the present work, recurrence rate 180 RR(t) and determinism – DET(t) measures based 181 on the analysis of diagonally oriented lines in the re-182 currence plot have been calculated [11]. Generally, re-183 currence rate RR(t) is the ratio of all recurrent states 184 (recurrence points) to all possible states and is there-185 fore the probability of the recurrence of a certain state. 186 The ratio of recurrence points forming diagonal struc-187 tures to all recurrence points is called the determin-188 189 ism DET(t). The larger values of RR(t) and DET(t)indicate the increase in regularity of investigated 190 dynamics. 191

## **3** Results and discussions

As an additional quantitative indication of phase syn-193 chronization between water level periodic variation and 194 seismic activity around large reservoir, observed in our 195 previous research [17], in the present research we have 196 calculated the phase diffusion coefficient, D. As fol-197 lows from Fig. 3, during the whole history of lake con-198 struction and exploitation, beginning from the territory 199 flooding (n = 1668 in Fig. 1(a)) and ending by regular 200 regime ( $n \approx 5000$ ), D indeed is minimal for the time 201 interval of periodic water level variation (Fig. 4). 20

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In the laboratory model of seismicity, acoustic emission during stick–slip experiments, it also was shown that phase diffusion coefficient D strictly decreases when acoustic emission time series are phase synchronized (see, e.g., Fig. 8 in [17]). 207

We explain the decrease of seismic energy release 208 shown in Fig. 1(c) which follows the period of RIS, by 209 phase synchronization of seismic activity with water 210 level variation (RISS). This seems to be a very im-211 portant example of purposeful man-made influence on 212 complex dynamics of seismic process. In this respect, it 213 should be mentioned that according to recent publica-214 tions data the dynamics of earthquake-related processes 215

**Fig. 3** Variation of phase diffusion coefficient of phase differences between daily released seismic energy and water level daily variations, calculated for consecutive sliding windows containing 365 events, shifted by 365 events (periodic forcing begins from  $n \approx 5000$ )





**Fig. 4** Variation of phase diffusion coefficient of phase differences between power of external sinusoidal forcing,  $P_s$  and power of acoustic emission of stick–slip events,  $P_a$  calculated for consecutive sliding windows containing 200 events, shifted by 200 events (periodic forcing of large enough amplitude begins from  $n \approx 2000$ )

in the earth crust are recognized as nonrandom, having 216 both low and/or high-dimensional nonlinear structures 217 [8, 13, 21, 24]. One of the characteristic properties of 218 processes in nonrandom systems, which are close to the 219 critical state, is their high sensitivity to initial conditions 220 as well as to relatively weak external influences. This 221 general property of complex systems acquires special 222 significance for practically unpredictable seismic pro-223 cesses. Indeed, insofar as we are not able to govern ini-224 tial conditions of lithospheric processes, even principal 225 possibility of controlling dynamics of seismic process 226 has immense scientific and practical importance (e.g., 227 to modify the release of accumulated seismic energy 228

via series of small or moderate earthquakes instead of one strong devastating event using the specific external impact). The way toward understanding such control mechanism passes through investigation of dynamics of seismic processes, when small external influence leads to phase synchronization. 234

It is known that nonlinear dynamical systems often 235 respond in a complicated way to such external influ-236 ences. One of the possible responses is synchroniza-237 tion since Huygens synchronization is understood as a 238 phenomenon when coupled nonlinear systems become 239 mutually adjusted. Presently, several types of synchro-240 nization are known, e.g., identical, generalized, phase 241 synchronization, etc. [4]. The phase synchronization 242 between water level periodic variation and seismic ac-243 tivity, observed in our previous and present researches, 244 is recognized as the weakest form of synchrony when 245 interacting nonlinear oscillators remain largely uncor-246 related [16, 20]. Generally, depending on the strength 247 of coupling, interacting systems may have different dy-248 namical features [16]. It is most important that con-249 trary to other forms of synchrony which lead to in-250 crease of order in behavior of synchronized system, 251 phase synchronization does not require strong cou-252 pling between involved processes. This, in turn, means 253 that the presence of order and character of changes 254 in dynamics of phase-synchronized system is not 255 obvious. 256

This is why we aimed to investigate the character of dynamical changes in seismic process when phase synchronization with water level periodic variation occurs. RQA, often used to detect changes in 260

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**Fig. 5** (a) RQA %DET of daily number of earthquakes calculated for consecutive non overlapping 1-year sliding windows (*circles*). Averaged results of RQA %DET for 20 shuffled (*asterisks*) and phase-randomized (*triangles*) surrogates of daily number of earthquakes in consecutive 1-year sliding windows; (b)

the dynamics of complex systems [9], is the most
convenient data analysis tool for this purpose. As
follows from our RQA results, when external influence on earth crust caused by water reservoir becomes

RQA %DET of magnitude (*black columns*) and waiting time (*grey columns*) sequences: (1) before impoundment, (2) during flooding and reservoir filling, and (3) periodic change of water level in reservoir

periodic, the extent of regularity of earthquake daily distribution (evaluated as %REC and %DET) essentially increases (see Fig. 5(a), bold line). This result was tested by comparing with the surrogate data. Averaged 266 267 268

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Fig. 6 RQA %DET calculated for consecutive nonoverlapping 400 data sliding windows of: (a) power of acoustic emission; (b) time intervals between consecutive maximal amplitudes of acoustic signals (*waiting times*). Averaged values for 20 shuffled time series are shown by triangles



results derived from RQA of 20 shuffled (asterisks) and 269 phase-randomized (triangles) surrogates (Fig. 5(a)), as-270 sure that the mentioned increase of regularity in earth-271 quakes distribution should not be an artifact. It is im-272 portant to mention that influence of increasing amount 273 of water and their subsequent periodic variation essen-274 tially affects also the character of earthquake's mag-275 nitude and temporal distribution (see Fig. 5(b)). Ex-276 tent of order in earthquakes temporal (black columns) 277 and magnitude (grey columns) distribution calculated 278

as value of %DET, substantially increases when the 279 reservoir forcing becomes periodic. Results of %DET 280 calculation of corresponding surrogates are always less 281 than 50% to original values (not shown here). It is inter-282 esting to mention, that dynamics of earthquakes tempo-283 ral and energetic distributions change even under water 284 level irregular variation though not so much as under 285 periodic variation. 286

The conclusions considered earlier on increase of 287 order in seismic process under water level periodic 288

variation using %DET measurements are confirmed by 289 calculation of other RQA measures (% REC, Entropy, 290 Laminarity). 291

As far as real-field seismic data sets are short and 292 incomplete, we carried out similar analysis on the 293 acoustic emission data sets, obtained on laboratory 294 spring-slider system under periodic electromagnetic 295 (EM) forcing, which simulates the periodical loading 296 by reservoir. Stick-slip experiments are considered as 297 a model of a natural seismic process [10, 21]. Time series of the emitted acoustic power during consecutive 299 cycles ( $2\pi$  periods) of the external 48 Hz periodic forc-300 ing of stick-slip process were analyzed as well as time 301 intervals between consecutive maximal amplitudes of 302 acoustic signals (waiting times). As is shown in Fig. 6 (circles), the extent of order increases both in ener-304 getic distribution as well as in temporal distribution 305 of acoustic emission when synchronization is achieved 306 (last window in Fig. 6). The averaged results of 20 307 surrogates shown by triangles confirm the conclusion 308 that observed changes are indeed related to ordering 309 in dynamics of acoustic emission under weak external 310 forcing. 311

#### 4 Conclusions 312

Dynamics of seismic process during RISS has been 313 investigated. Data sets of daily water level variation 314 and released seismic energy as well as waiting time 315 and magnitude sequences were analyzed. As a model 316 of natural seismicity, the laboratory stick-slip acous-317 tic emission data were also analyzed. Methods of phase diffusion coefficient calculation and RQA were 319 used. 320

Based on the results of investigation carried out 321 both on field and experimental time series, we con-322 clude that the order in dynamics of earthquake's 323 daily occurrence, as well as in earthquake's tempo-324 ral and energetic distributions increases when water 325 level variation becomes periodic. Laboratory stick-slip 326 acoustic emission data confirms results of field data 327 analysis. 328

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## References

- 1. Assumpção, M., Marza, V.I., Barros, L.V., Chimpliganond, C.N., Soares, J.E.P., Carvalho, J.M., Caixeta, D., Amorim, A., Cabral, E.: Reservoir induced seismicity in Brazil. Pure Appl. Geophys. 159, 597-617 (2002)
- 2. Belaire-Franch, J., Contreras, J.D., Tordera-Lledo, L.: Assessing nonlinear structures in real exchange rates using re-341 currence plot strategies. Physica D 171, 249-264 (2002)
- 3. Bowman, D.D., Ouillon, G., Sammis, C.G., Sornette, A., Sornette, D.: An observational test of the critical earthquake concept. J. Geophys. Res. 103, 24359-24372 (1998)
- 4. Calvo, O., Chialvo, D., Eguíluz, V., Mirasso, C., Toral. R.: Anticipated synchronization: a metaphorical linear view. Chaos 14, 7-13 (2004)
- 5. Chelidze, T., Lursmanashvili, O.: Electromagnetic and mechanical control of slip: laboratory experiments with slider system. Nonlinear Process. Geophys. 20, 1-8 (2003)
- 6. Chelidze, T., Matcharashvili, T., Gogiashvili, Lursmanashvili, O., Devidze, M.: Phase synchronization of slip in laboratory slider. Nonlinear Process. Geophys. 12, 1–8 (2005)
- 7. Eckmann, J.P., Kamphorst S., Ruelle, D.: Recurrence plots of dynamical systems. Europhys. Lett. 4(9), 973-977 (1987) 357
- 8. Goltz, C.: Fractal and Chaotic Properties of Earthquakes. Springer, Berlin (1998)
- 9. Iwanski, J., Bradley, E.: Recurrence plots of experimental 360 data: to embed or not to embed? Chaos 8(4), 861-871 (1998) 361
- 10. Johansen. A., Sornette, D.: Acoustic radiation controls dynamic friction: evidence from a Spring-Block Experiment. Phys. Rev. Lett. 82, 5152-5155 (1999)
- 11. Marwan, N, Wessel, N., Meyerfeldt, U., Schirdewan, A., 365 Kurths, J.: Recurrence-plot-based measures of complexity 366 and their application to heart rate variability data. Phys. Rev. 367 E 66, 026702.1-026702.8 (2002) 368
- 12. Marwan, M.: Encounters with neighborhood. PhD Thesis 369 (2003)3742
- Matcharashvili, T., Chelidze, T., Javakhishvili, Z.: Nonlin-371 ear analysis of magnitude and interevent time interval se-372 quences for earthquakes of Caucasian region. Nonlinear Pro-373 cess. Geophys. 7, 9-19 (2000)
- 14. McAllister, R., Uchida, A., Meucci, Roy, R.: Generalized synchronization of chaos: experiments on a two-mode microchip laser with optoelectronic feedback. Physica D 195, 244-262 (2004)
- 15. Nascimento, A.F., Cowie, P.A., Lunn, R.J., Pearce, G.: 379 Spatio-temporal evolution of induced seismicity at Açu 380 reservoir, NE Brazil. Geophys. J. Int. 158, 1041-1052 381 (2004)
- 16. Pazo, D., Zaks, M.A., Kurths, J.: Role of unstable periodic 383 orbits in phase and lag synchronization between coupled 384 chaotic oscillators. Chaos 13, 309-318 (2003) 385
- 17. Peinke, J., Matcharashvili, T., Chelidze, T., Gogiashvili, J., 386 Nawroth, A., Lursmanashvili, O., Javakhishvili, Z.: Influ-387 ence of periodic variations in water level on regional seismic 388 activity around a large reservoir: field and laboratory model. 389 Phys. Earth Planet. Inter. 156(1-2), 130-142 (2006) 390
- 18. Pikovsky, A., Rosenblum, M.G., Kurth. J.: Synchronization: 391 universal Concept in Nonlinear Science. Cambridge Univer-392 sity Press, Cambridge, MA (2003) 393

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- 19. Postnov, D.E., Sosnovtseva, O.V., Mosekilde, E., Holstein-
- Rathlou, N.-H.: Synchronization of tubular pressure oscillations in interacting nephrons. Chaos Solitons Fractals 15(2),
   343–369 (2003)

S43–S09 (2003)
 Rosenblum, M., Pikovsky, A., Kurth, J.: Phase synchroniza-

- tion of chaotic oscillators. Phys. Rev. Lett. 76, 1804–1807
   (1996)
- 401 21. Rundle, J., Turcotte, D., Klein, W. (eds): GeoComplexity and
   402 the Physics of Earthquakes. American Geophysical Union,
- 403 Washington, DC (2000)
- 404 22. Simpson, D.W.: Triggered earthquakes. Annu. Rev. Earth Planet. Sci. 14, 21–42 (1986)
- 23. Simpson, D., Leith, W., Scholz, C.: Two types of reservoirinduced seismicity. Bull. Seismol. Soc. Am. 78, 2025–2040 406 (1988) 407
- 24. Smirnov, V.B.: Fractal properties of seismicity of Caucasus. J. Earthq. Predict. Res. 4, 31–45 (1995) 409
- 25. Talwani, P.: On nature of reservoir-induced seismicity. Pure Appl. Geophys. **150**, 473–492 (1997) 411
- 26. Trifu, C.I. (ed.): The mechanism of induced seismicity, special volume. Pure Appl. Geophys. 159 (2002)
   4A3
- 27. Zbilut, J.P., Webber, C.L. Jr.: Embeddings and delays as derived from quantification of recurrence plots. Phys. Lett. A 171, 199–203 (1992)

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